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A 2.5-D Diffraction Tomography Inversion Scheme for Ground Penetrating Radar

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Abstract — A new 2.5-D inversion scheme is derived for ground penetrating radar (GPR) that applies to a monostatic fixed-offset measurement configuration. The inversion scheme, which is based upon the first Born approximation and the pseudo-inverse operator, takes rigorously into account the planar air-soil interface, the loss in the soil, and the characteristics of the antennas.

1. Introduction

Several inversion schemes, based upon the first Born approximation and the concept of diffraction tomography (DT) [1], have been derived for monostatic ground penetrating radar (GPR) configurations [2–5]. In practical situations the GPR is usually situated upon a planar interface that separates air from soil and the soil is usually lossy. Therefore, it is important – as illustrated in [5] – to incorporate in the inversion the presence of the planar interface and the fact that the soil has loss. However, the inversion schemes in [2, 5] do not take into account the loss in the soil and only the one presented in [5] includes the planar air-soil interface. The present paper derives, to the knowledge of the author, the first DT inversion scheme that accounts for both the planar interface and the lossy soil. The starting point is the forward model of [5] which is based upon the first Born approximation, the dyadic Green function for the planar air-soil interface, and an asymptotic approximation valid when the object is located deep (a few wavelengths) in the soil. However, instead of inverting this forward model using the inverse Fourier transform, as done in [5] – and thus neglecting loss in the soil – the pseudo-inverse operator of [3, 4] is used.

2. The Forward Model

The GPR configuration is shown in Figure 1 in which a planar interface separates air from soil. A Cartesian xyz coordinate system is introduced such that the xy plane coincides with the interface and such that $z > 0$ is air. An object, which is assumed infinitely long in the \hat{x} direction, is buried in the soil. The propagation constant of air is $k_0(\omega) = \omega\sqrt{\mu_0\epsilon_0}$ and that of soil is $k_1(\omega) = \omega\sqrt{\mu_0(\epsilon_1 + i\sigma_1/\omega)}$ (time factor $\exp(-i\omega t)$). The position of the receiving antenna is described by $\mathbf{r}_r = \mathbf{R}_r + \hat{z}z_r$, and that of the transmitting antenna is $\mathbf{r}_t = \mathbf{r}_r + \mathbf{r}_\Delta$ with the fixed offset $\mathbf{r}_\Delta = \mathbf{R}_\Delta + \hat{z}z_\Delta$. It is possible to derive a forward model that holds for arbitrary antennas [5]. However, for the sake of simplicity, it is here assumed that the antennas can be accurately modeled by Hertzian dipoles. It is also assumed that the contrast in conductivity $\Delta\sigma(y, z) = \sigma(y, z) - \sigma_1$ is much less than the contrast in permittivity, i.e., $\Delta\sigma(y, z) \ll \omega\Delta\epsilon(y, z)$ over the frequency band of interest $\omega_{\min} < \omega < \omega_{\max}$. Then $\Delta\epsilon(y, z)$ can be related to the output of

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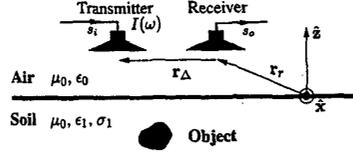


Figure 1: The fixed-offset monostatic GPR configuration.

the receiving antenna s_o as [5, (11)]

$$\bar{s}_o(k_y, z_r, \omega) = (L\Delta\epsilon_1)(k_y, z_r, \omega) = -i\omega D(k_y, z_r, \omega) \int_{-\infty}^{\infty} dy' \int_{-\infty}^0 dz' \cdot \exp(-ik_y y') \exp(-i\sqrt{4k_1^2(\omega) - |k_y|^2} z') \Delta\epsilon_1(y', z') \quad (1)$$

where $\Delta\epsilon_1 = \Delta\epsilon/z'$ and the Fourier transform \bar{s}_o of the output of the receiving antenna is obtained from $\bar{s}_o(k_y, z_r, \omega) = \int_{-\infty}^{\infty} dy_r s_o(y_r, z_r, \omega) \exp(-ik_y y_r)$. The linear operator $L : U \rightarrow V$ maps the space U onto the space V . U is the space of square integrable functions of position (y', z') confined within $z' < 0$. V is the space of square integrable functions defined on $\{(k_y, \omega) | \omega_{\min} < \omega < \omega_{\max} \wedge |k_y| < 2\text{Re}k_1(\omega)\}$. Since $\bar{s}_o = 0$ for $|k_y| > 2\text{Re}k_1(\omega)$ no evanescent plane waves in the soil are considered. Moreover, when the Hertzian dipoles are \hat{x} -directed and have the same x and z coordinates ($x_{\Delta} = z_{\Delta} = 0$), the function D in (1) is

$$D(k_y, z_r, \omega) = \frac{i\omega^2 \mu_0^2 I(\omega) (4k_1^2(\omega) - k_y^2)}{4\pi k_1(\omega) \left(\sqrt{4k_0^2(\omega) - k_y^2} + \sqrt{4k_1^2(\omega) - k_y^2} \right)^2} \cdot \exp\left(i \left[\sqrt{4k_0^2(\omega) - k_y^2} z_r + \frac{1}{2} k_y y_{\Delta} \right] \right) \quad (2)$$

where $I(\omega)$ is the impressed current of the transmitting Hertzian dipole.

3. Inversion

The forward model (1) is now inverted using the (Tikhonov-regularized) pseudo-inverse operator [6, p. 88]

$$\Delta\epsilon = z' L^{\dagger} (L L^{\dagger} + \lambda^2 I)^{-1} \bar{s}_o \quad (3)$$

where the adjoint operator L^{\dagger} , defined by $\langle \bar{s}_o, L\Delta\epsilon_1 \rangle_V = \langle L^{\dagger} \bar{s}_o, \Delta\epsilon_1 \rangle_U$ with the usual definition of the inner products, is

$$(L^{\dagger} \bar{s}_o)(y, z) = \int_{\omega_{\min}}^{\omega_{\max}} d\omega i\omega \int_{|k_y| < 2\text{Re}k_1(\omega)} dk_y D^*(k_y, z_r, \omega) \cdot \exp(ik_y y) \exp(i\sqrt{4k_1^2(\omega) - k_y^2} z) \bar{s}_o(k_y, z_r, \omega). \quad (4)$$

Herein, * denotes the complex conjugate. The quantity λ in (3) is a regularization parameter. However, in the numerical example in Section 4 a situation is considered which does not require regularization, i.e., $\lambda = 0$. Now, the *filtered data* \tilde{s}_o^f is introduced as the solution to

$$(LL^\dagger + \lambda^2 I)\tilde{s}_o^f = \tilde{s}_o. \quad (5)$$

Using this definition along with (3), the contrast in permittivity is obtained from

$$\Delta\epsilon = z' L^\dagger \tilde{s}_o^f. \quad (6)$$

Hence, by solving (3) using the solution steps (5) and (6), the data are first filtered and then backpropagated to obtain the sought-for function $\Delta\epsilon$. Observe that the term LL^\dagger in the filtering step (5) can be explicitly expressed as

$$(LL^\dagger \tilde{s}_o^f)(k_y, \omega) = 2\pi i \omega D(k_y, z_r, \omega) \int_{\max(\omega_{\min}, \omega_m(k_y))}^{\omega_{\max}} d\omega' \omega' \frac{D^*(k_y, z_r, \omega') \tilde{s}_o^f(k_y, \omega')}{\sqrt{4k_y^2(\omega) - |k_y|^2} - \sqrt{4k_y^2(\omega') - |k_y|^2}} \quad (7)$$

where $\omega_m(k_y)$ satisfies the relation $2\text{Re}k_1(\omega_m(k_y)) = |k_y|$. Hence, when inserting (7) into (5), an integral equation is obtained for determination of the filtered data \tilde{s}_o^f for each value of k_y . This integral equation is solved numerically using pulse expansion and point matching. To this end, assume that the radar data \tilde{s}_o is available at the equidistant frequencies $\omega_p = (p-1)\Delta\omega + \omega_{\min}$, $\Delta\omega = (\omega_{\max} - \omega_{\min})/(N_\omega - 1)$, $p = 1, \dots, N_\omega$. With $q(k_y)$ being the lowest positive integer satisfying $\omega_m(k_y) < \omega_p$, the filtered data is expanded as $\tilde{s}_o^f(k_y, \omega) = \sum_{p=q(k_y)}^{N_\omega} s_{op}^f(k_y) u_p(\omega)$ where the first $q(k_y) - 1$ values of $s_{op}^f(k_y)$ equal zero. The pulse expansion functions $u_p(\omega)$ are

$$u_p(\omega) = \begin{cases} 1 & \text{for } \omega_{\min} < \omega < \omega_{\min} + \Delta\omega/2 \text{ and } p = 1 \\ & \text{or for } \omega_{\max} - \Delta\omega/2 < \omega < \omega_{\max} \text{ and } p = N_\omega \\ & \text{or for } |\omega - \omega_p| < \Delta\omega/2 \text{ and } 1 < p < N_\omega. \\ 0 & \text{otherwise.} \end{cases} \quad (8)$$

The resulting matrix equation for obtaining $\tilde{s}_o^f(k_y)$ is easily derived by inserting the expansion for $\tilde{s}_o^f(k_y, \omega)$ into (7) and using (5) with point matching.

4. Numerical Example

The inversion scheme of Section 3 is now tested on synthetic GPR data. Figure 2 shows a dielectric pipe with a diameter of 15cm and electromagnetic properties $(\epsilon_{\text{pipe}}, \sigma_{\text{pipe}}) = (8.1\epsilon_0, 0.01\text{S/m})$ located 1m below the interface. The soil has $(\epsilon_1, \sigma_1) = (8\epsilon_0, 0.01\text{S/m})$. The synthetic GPR data is calculated from an eigenfunction expansion. It is assumed that the radar uses 60 frequencies equally spaced in the range $20 \text{ MHz} < f < 1.3 \text{ GHz}$. Figure 3 shows the image of $\Delta\epsilon(y, z)/\epsilon_0$ obtained from (5) with $\lambda = 0$ and from (6). The image is of high quality and it approximates well the correct value of $\Delta\epsilon = 0.1\epsilon_0$.

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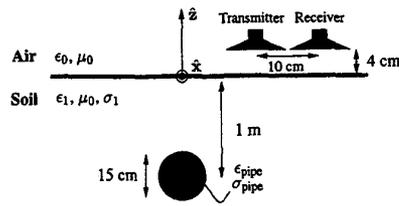


Figure 2: The configuration involving a pipe with $\Delta\epsilon = 0.1\epsilon_0$.

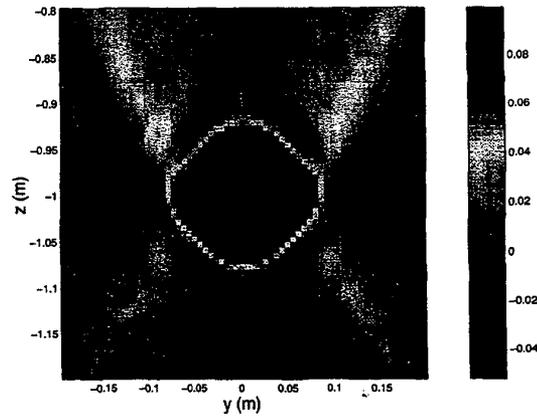


Figure 3: The image of $\Delta\epsilon(y, z)/\epsilon_0$.

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